Remote Sensing of Glaciers Chapman Glacier, Ellesmere Island, Nunavut – ASTER 2000



Landsat Images (since 1972 / 1984) Most glaciers are remote

Note mark of Little Ice Age ~ 1850 LIA: ~ 1450-1850



Note late lying snow cover





Glacier facies are natural zones of distinct variations of snow and ice which are formed as a result of the evolution of precipitated snow to ice, the cyclic process of ablation, refreezing, and eventually its melt.

Accumulation

Ablation

Hudson Bay Glacier 1912-2000



Robson Glacier 1911-2010



Robson Glacier 1911 - Byron Harmon



Robson Glacier 2010 - Robin Draper



The spectral curve explains why glaciers look blue-green on a 5-4-3 composite (why?) .. and enables distinguishing snow/ice from clouds compared to a normal composite.. (why?)



Mid-IR/Near-IR-Red



Red-Green-Blue

http://asterweb.jpl.nasa.gov/gallery-detail.asp?name=Aletsch



1. Image classification - supervised

Accumulation area (snow); Ablation (ice) ... and firn (wet snow)



Image processing

1. Classification: 2000 areas produced by supervised classification

Edziza: extents from NTDB 1966, BC TRIM 1985, Landsat 2000

Training on ablation / accumulation areas



Unsupervised classification: McBride OLI image including Kristi Glacier (SW corner)



Image classification - Unsupervised



These orange-pink clusters, not the brown one (forefield) – why so many - 6?

2. Normalised Difference Snow Index (NDSI): green-SWIR / Green + SWIR

NDSI (TM) = (2-5)/(2+5)

NDSI (OLI) = (3-6)/(3+6)

Method: use as threshold or input in classification

Note: its impossible to distinguish between snow covering glaciers and late lying snow on land except by size (sieve) and perhaps modelling from location

3. Ratio image - thresholding

.... NIR/MIR band ratio TM 4 / 5 (snow/ice >1.0) Red/MIR TM 3/5 (snow/ice > 2.0) ... 'better' for shadow areas



Snow and ice: very high in visible, very low in Mid-Infrared ('SWIR') **Ratio - Visible (Red) to SWIR** captures snow/ice 'almost' exclusively - Some issues with silt-laden water, shadowed glaciers and debris cover

Landsat 1992

2 km



Carstenz Icefield, Indonesia, 4 S - no shadows, 5000m elevation

Red/SWIR ratio



Threshold value 2.0



Convert bitmap to polygon



Vector smoothing



Overlay of 2017 polygons on 1992 image showing ice loss



The cordilleran glaciers of western Canada





with threshold >2.0 \sim 15,000 glaciers covering \sim 25,000 km² ('we' = Tobias Bolch !)

Change per year (%) More glacier mass in the Coast Mountains



Glacier retreat 1985 - 2015 at Mount Robson





Image data: Landsat B (NASA) Outline data: T. Bolch (2005) and TRIM (BC Govt.) (1985) Projection: WGS 84 UTM Zone 10N Cartography: Sophia Goertz Date: 10.09.2018

Svalbard subset overview (bands 8 4 3) (Frank Paul)



Resulting corrected outlines (Frank Paul)



Ongoing Challenges

1. mixed pixels \rightarrow lower threshold

2. shadows → lower threshold

... view in visible bands or Panchromatic

- 3. Misclassified lakes → higher threshold
- 4. Debris Cover DEM? Thermal?



Later in the year: less snow, more shadow





opernicus

Approach: map with July scene, correct with September

Svalbard: 80 N







Remote Sensing of Glaciers

Image processing can be used to map:

a. Glacier extents (compare with map vectors)

- b. Surface characteristics (e.g. accumulation-ablation)
- c. Animation image series

d. Glacier movement /velocity

e. Elevation / volume change with DEMS (see next lecture)



d. Glacier velocity

Klinaklini Glacier

Annual movement ranges 30 - 500m (mostly summer) ~ 1m per day (summer)

Length of vector proportional to change between sequential Images Oct 2001/Sep 2002

ENVI COSI-CORR software

SPOT high resolution imagery 2.5m

Scud Glacier (2002)

Scud Glacier (2003)

0.5 km

e. DEMs - more next lecture Image processing can be used to map and measure : Elevation change / Volume loss from DEMs

